

# The Mapping of Landscapes, Geology and Soils of Bedfordshire & Cambridgeshire



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We use the words landscape and landform to refer to natural (not man-made) features visible on the surface of the Earth. The word landscape is used for larger areas of varied topography, whereas the word landform is used for smaller features that occur within landscapes, and can often be readily seen to have resulted from the local action of particular surface processes (for example, slope failure or river erosion). In this study we present a group of regional landscape maps (Maps 1-4), and a group of more local maps (Maps 5-8) on a larger scale, where landforms are more easily distinguished. Our work has been part-funded by Natural England and facilitated by Geo-East, the East of England Geodiversity Partnership.

Landscapes and their component landforms have formed during the long-term geological history of an area, and may have been influenced by many factors. These include, a) the materials present just below the Earth's surface, b) movements of the Earth's land or sea surface, and c) the action of ice, rain, wind and living organisms. This study has been concerned particularly with ways of analysing and presenting topographical information, so that members of the general public can gain new insights into the stories that have resulted in their landscapes and landforms.

Maps are routinely used to represent landscapes and landforms, and we have wanted to experiment with the ways that computer-based Geographical Information System (GIS) software, using the digital regional datasets now available, has transformed our ability to provide vivid representations of our surroundings.

The approach we have used has been to derive an elevation model from the OS Opendata Panorama dataset, vertically exaggerated by a factor of 20, and then apply colour and shading to illustrate topographic relief (Maps 1 and 5). We have then added overlays to compare the topography with bedrock geology (Maps 2 and 6), surface geology (Maps 3 and 7), and soil distribution (Maps 4 and 8). The geological datasets are derived from the BGS DiGMap50 dataset, which have been merged and rationalised to provide a consistent coverage across the whole region. The soils datasets are the NSRI National Soil Map of England and Wales (NATMAPvector) which describes soils distribution in the landscape. The GIS processing was done in ArcGIS.

The area covered in this study is Bedfordshire, Cambridgeshire, and the Borough of Milton Keynes (part of Buckinghamshire). As an example of the potential for featuring landforms more clearly at the local scale, we have focused on the area of Marston Vale, southwest of Bedford. This 25 x 20 km area is highlighted on Map 1, and forms the basis for Maps 5-8.

Further information on these topics is covered on the following websites:

Bedfordshire Geology Group  
[www.bedfordshiregeologygroup.org.uk](http://www.bedfordshiregeologygroup.org.uk)

Cambridgeshire Geology Club  
[www.cambridgeshiregeologyclub.org.uk](http://www.cambridgeshiregeologyclub.org.uk)

National Soil Resources Institute, Cranfield University  
[www.landis.org.uk](http://www.landis.org.uk)

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River Nene draining eastwards into the Fens

Holme Fen (Cams), the lowest land point in Britain at 2.75m below sea level

The man-made drainage is apparent in the Fens

Confluence of the River Great Ouse and Alconbury Brook

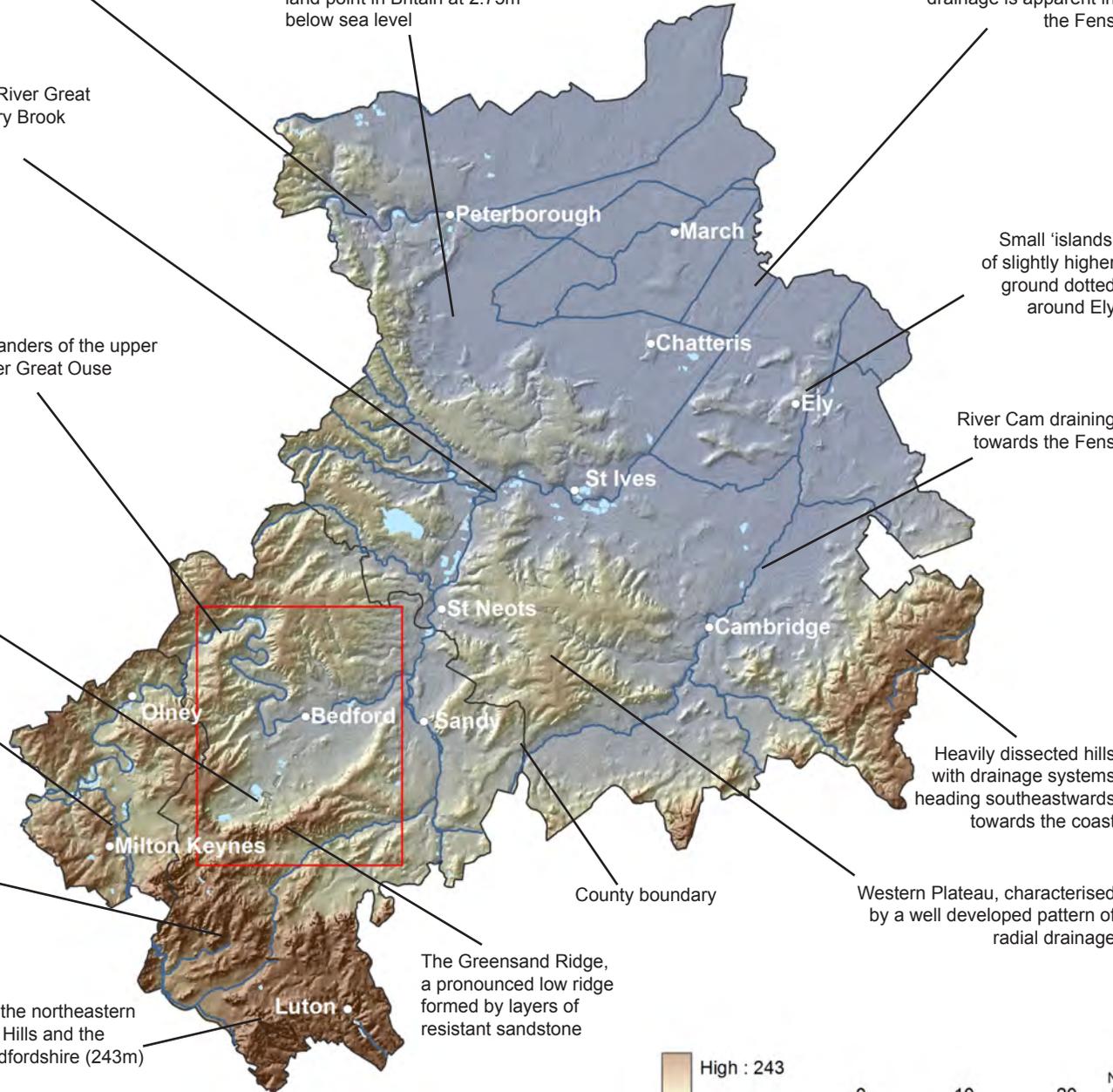
Deeply incised meanders of the upper reaches of the River Great Ouse

Marston Vale

River Ouzel and the Grand Union Canal

Heavily dissected high ground flanked by low-lying clay vales

Dunstable Downs, the northeastern part of the Chiltern Hills and the highest point in Bedfordshire (243m)



# Map 1. Topography

The area is generally low-lying with 50% of the land lower than 30 m. It is drained by two major rivers, the Great Ouse and Nene, and their tributaries. Rivers and water courses are shown on this and all subsequent maps in blue.

The arc of low hills between Peterborough and Ely define the inland extent of the Fens, a remarkably flat and extensive plain caused by numerous inundations of sea water during the last few thousand years. The Fens are now drained by a series of engineered water courses that have changed the natural river systems completely.

There are several large, flat-topped plateaux in south Cambridgeshire (e.g. the Western Plateau) and north Bedfordshire (e.g. flanking the Great Ouse). They were commonly used in the past as locations for military airfields.

There are many mid-height, linear ridges in Bedfordshire that trend NE-SW. Some result from hard layers within the bedrock geology (see Map 2), but most are formed by flanking valleys cut by streams flowing northeastward towards the lowest part of the region.

High ground to the southeast of Cambridge is formed by the same relatively hard and resistant rock (Chalk) that occurs around Luton at similar elevations. The Chalk ridge continues southwestwards, with increasingly strong definition, as the Chiltern Hills.

The red rectangle highlights the area around Marston Vale (see Maps 5-8).

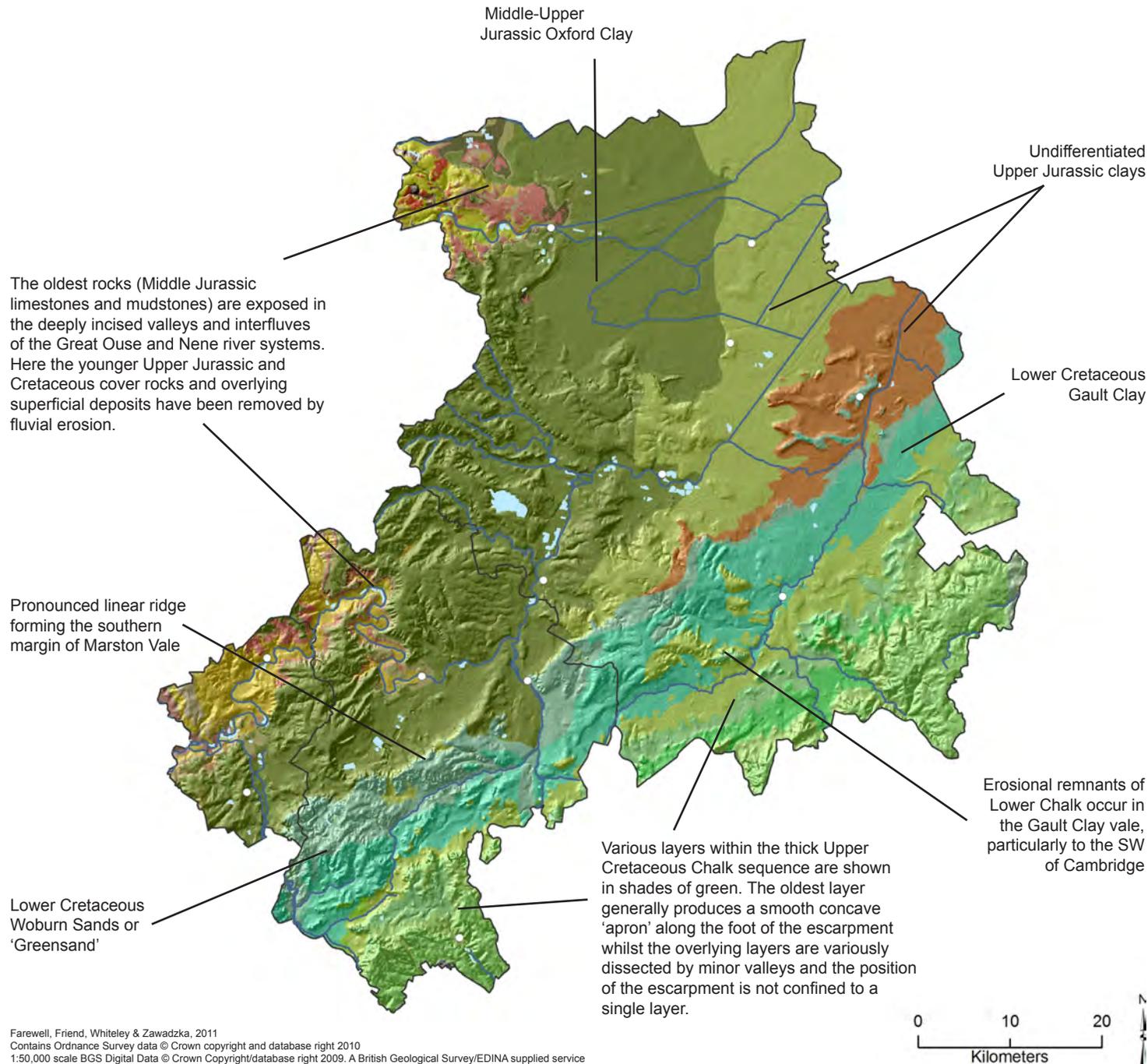
## Map 2. Bedrock Geology

This map shows how the rocks would appear at the land surface if the overlying superficial deposits and soil were removed. Described as bedrock, these rocks are all of sedimentary origin but they vary considerably in terms of their age (from about 185 million to 85 million years old) and their type (mainly mudstones, limestones and sandstones).

The distribution of bedrock geology is not coincident with most of the landscapes. For example, the extensive tract of Middle-Upper Jurassic clays underlies both the flat-lying Cambridgeshire Fens and the heavily dissected hills of north Bedfordshire. A similar pattern is observed along the extent of the Gault Clay outcrop. It is noticeable that the highest tracts of ground occur in the south-west across the complete range of rock types.

Lower Cretaceous sandstones in Bedfordshire are relatively resistant and produce an intermittent escarpment and broad, dissected dip slope known locally as the Greensand Ridge. This feature is particularly clear around Sandy where it is cut by the River Ivel.

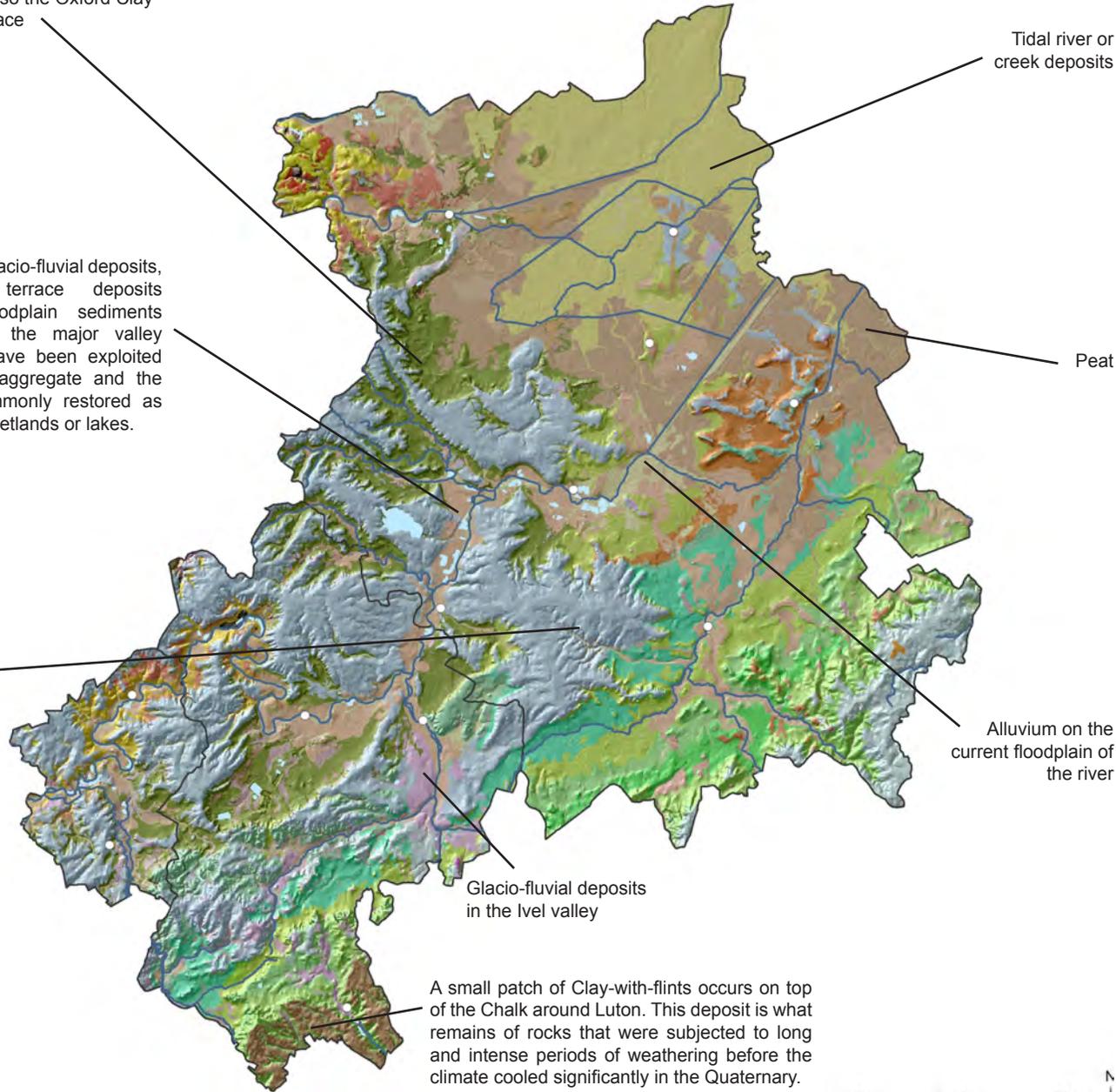
The scarp and vale landscape that characterises the southern half of the study area developed initially through differential erosion under sub-tropical and warm temperate climates during the Tertiary. When a major ice sheet advanced and retreated across the area some 450-400,000 years ago it modified that landscape enormously, initially grinding much of it away, then blanketing it with considerable thicknesses of glacial deposits. Since then, alternating cycles of cool temperate and very cold climates have worn away much of the glacial material and revealed some of the underlying bedrock again.



Area where the superficial deposits are thin or absent, so the Oxford Clay is close to the surface

A wide variety of glacio-fluvial deposits, younger river terrace deposits and modern floodplain sediments occur throughout the major valley systems. Many have been exploited commercially for aggregate and the workings are commonly restored as agricultural land, wetlands or lakes.

Anglian till



Tidal river or creek deposits

Peat

Alluvium on the current floodplain of the river

Glacio-fluvial deposits in the Ivel valley

A small patch of Clay-with-flints occurs on top of the Chalk around Luton. This deposit is what remains of rocks that were subjected to long and intense periods of weathering before the climate cooled significantly in the Quaternary.

### Map 3. Surface Geology

This map superimposes the extensive 'blanket' of glacial and post-glacial deposits on the bedrock geology. Known as superficial deposits, they formed during the Quaternary, a period from about two million years ago to the present day that is characterised by alternating cycles of cold and temperate climate.

The most obvious feature of the superficial deposits are the widespread cappings of glacial till, a clay-rich deposit that includes lumps of chalk, sandstone, limestone and flint. The till is relatively resistant and forms high ground across most of the area. In parts of north Bedfordshire and on the Western Plateau it is >40 m thick and it formed as a major ice sheet (Anglian) pushed southwards towards the Chilterns, churning up the underlying bedrock and re-depositing it when the ice sheet subsequently melted about 400,000 years ago.

Clay-with-flints are limited in extent, suggesting that only the highest hill tops in Bedfordshire were beyond and above the erosive limit of the Anglian ice sheet. Today they form a resistant and impermeable capping on the interfluvial valleys cut deep into the Chalk dip slope.

The Fens provide Britain's largest and flattest landscape. At elevations close to sea level they have been repeatedly inundated by sea water and progressively buried by fine-grained deposits that effectively suppress any underlying topography. Further inland, in a broad arc from Peterborough to Ely, peat has formed as a result of the waterlogged and swampy conditions that existed behind the transient coastal zone.

In these flat areas, the level of the groundwater is generally controlled by ditches and pumps

These deep stoneless clayey soils have developed over marine alluvium and are often quite acid

The marine alluvium and Fen peats give rise to soils which are seasonally wet, deep clays. These are sometimes calcareous in places.

Shallow well drained brashy calcareous fine loamy soils have developed over limestone

Deep loamy soils have formed over the chalky till and are used for growing both cereals and field vegetables

The River Great Ouse cuts down through the till and underlying mudstones into the Jurassic limestones below. This gives rise to typical brown calcareous earths. Alluvium with a clayey, calcareous subsoil is found on the floodplains.

The cherty, flinty and sometimes chalky Clay-with-flints and sandstones give rise to deep loam to clay soils

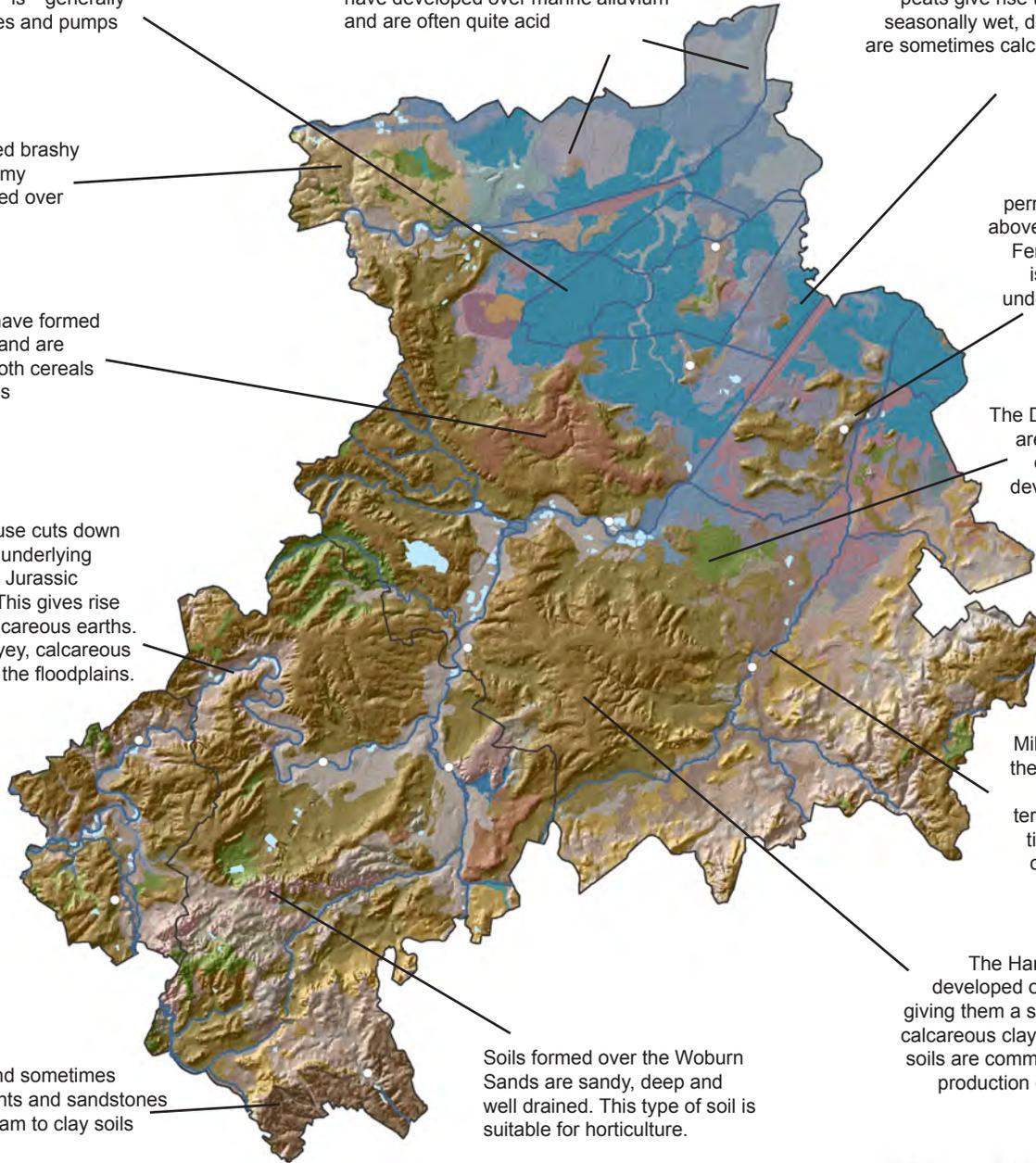
Soils formed over the Woburn Sands are sandy, deep and well drained. This type of soil is suitable for horticulture.

The deep, slowly permeable clays rise above the surrounding Fens. The elevation is provided by the underlying chalky till.

The Denchworth soils are seasonally wet, deep clays which develop on Jurassic mudstones.

The permeable Milton soils flanking the River Cam have formed on river terraces and chalky till and give rise to complex local soil patterns

The Hanslope soils have developed on the chalky tills, giving them a slowly permeable, calcareous clayey nature. These soils are commonly used for the production of winter cereals.



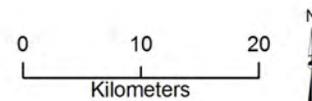
## Map 4. Soils

This map shows Soil Associations, which are groupings of soils commonly found in association with each other in the landscape. These are named after type localities.

Soil performs a number of key functions that are essential for our existence. It is well recognized that soil is the growing medium for plants and crops, but it is equally important to note that the soil itself is home to a complex and active ecosystem with bacteria, fungi, protozoa, nematodes, insects and worms driving the cycling of nutrients, regulating plant communities and helping repair soil structure. Soil filters, cleans and breaks down pollutants, thus helping to ensure good water quality. The soil also captures approximately 20% of the anthropogenic carbon we create each year.

Soil development is a function of climate, topography, vegetation, organisms and humans, geological parent material and time. Across this region, the two most important natural factors leading to the observed diversity in soils are the differences in the underlying geological parent material and the relationship between soil and topography.

The soils have developed since the Anglian ice sheet retreated from the area about 400,000 years ago. As such, they have formed on a variety of glacial and post-glacial deposits or directly on the underlying bedrock. The topography interacts with climate and vegetation, directing soil formation and controlling soil-water budgets. For example, steep slopes promote surface run-off and reduced leaching, whilst valley floors are often waterlogged. These interactions continue to affect the ongoing development of the soil.



Northeasterly Great Ouse meander reach (c.50 – 40 m river elevation)

Southerly Great Ouse meander reach (c.40 – 25 m river elevation)

Northeasterly trending large ridges with unfluted margins

Brogborough 'saddle', elevation (c.92 m)

Highest elevation (c.130 m)

Fluted basin margins

Lowest elevation (c.20 m)

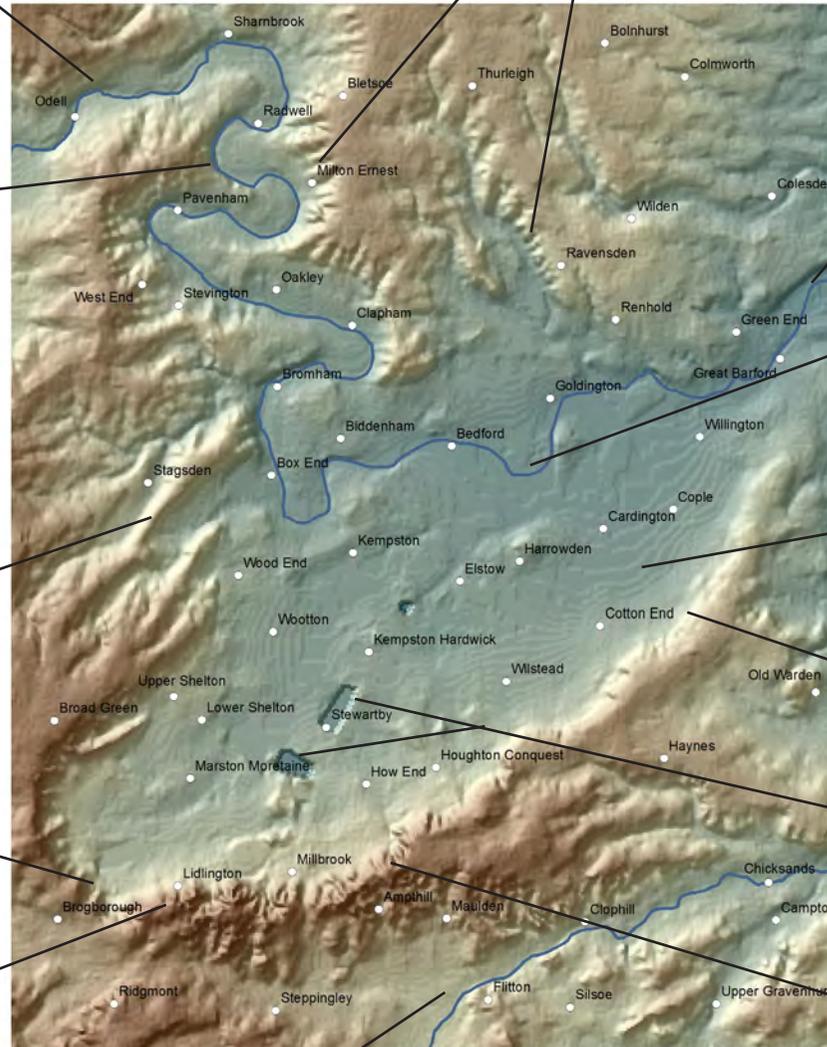
Easterly Great Ouse reach through Bedford (c.25 – 20 m river elevation) and towards the sea

Apparent 'terrace' effect is an artefact of the digital elevation model

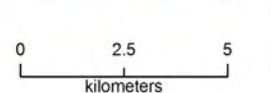
Unfluted margin on northeasterly trending ridge

Disused clay pits that provided the raw material for the 20th century brick-making industry

Fluted basin margin



Valley of the River Flit



## Map 5. Topography

Marston Vale (located on Map 1) is a distinctive low-lying basin, extending some 25 km in a southwest to northeast direction. Apart from man-made clay and gravel pits, the vale has a generally flat floor with elevations averaging about 35 m above sea-level. The floor is surrounded by higher ground except at the entry and exit points of the River Great Ouse.

The Ouse is the largest river system in East Anglia. It enters the map area from the northwest with two distinctive reaches of strongly meandering incised valley upstream of Bedford, and then continues flowing eastwards through more poorly defined topography that forms the northeastern end of Marston Vale. South of the Great Ouse, the floor of Marston Vale currently lacks a clearly defined local drainage system of stream channels or valleys.

Many, but not all, of the slopes in this map area provide examples of a distinctive local landform of ridges and small valleys running more or less at right angles to the general trend of the larger slope, and usually tens to a few hundred metres across. In the discussion that follows we find it useful to refer to these landforms as 'flutes'.

Middle Jurassic (Great Oolite Group) Blisworth Limestone Formation

Middle Jurassic (Great Oolite Group), Cornbrash Formation (limestone)

Middle Jurassic Kellaways Formation (sandstone and mudstone)

Minor valleys or 'flutes' coinciding with the extent of the Woburn Sands Formation along the southern margin of the vale

Lower Cretaceous Woburn Sands Formation (sandstone)

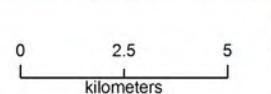
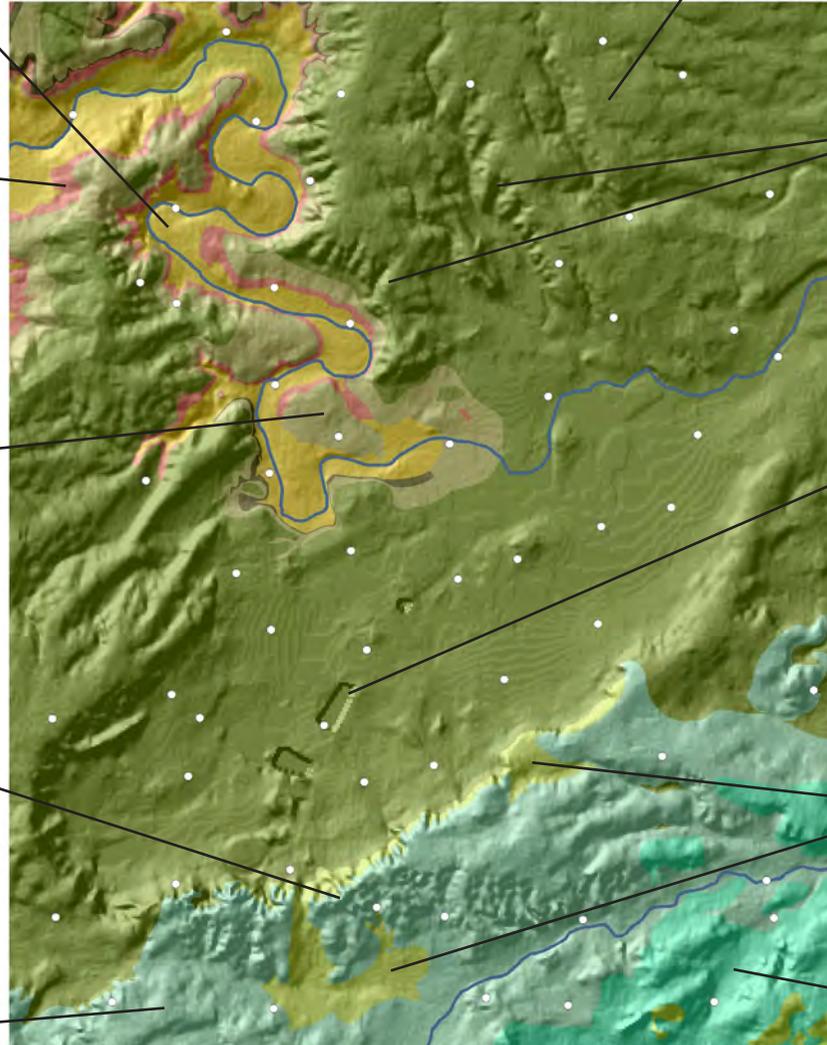
Middle to Upper Jurassic Oxford Clay Formation (mudstone)

Minor valleys or 'flutes' are orthogonal to major N-S trending valleys

The oldest part of the Oxford Clay Formation (Peterborough Member) floors most of the vale and was exploited for brick-making

Upper Jurassic West Walton and Amphill Formations (mudstone)

Lower Cretaceous Gault Clay Formation (mudstone)



## Map 6. Bedrock Geology

The oldest bedrock visible has been exposed in the northwest of the map area by the erosional down-cutting of the River Great Ouse. It consists largely of limestones, though mudstones are also important at some levels, and these sediments are classified as part of the Great Oolite Group.

Above these mainly Middle Jurassic rocks, the rest of the Jurassic bedrock succession is classed as Ancholme Group, which includes the Kellaways, Oxford Clay, West Walton & Amphill Clay formations. These are mainly mudstones, some of which have a relatively high organic carbon content that made them particularly useful in the brick-making industry around Stewartby.

Lower Cretaceous bedrock is present in the near surface of the southern part of the Marston Vale area. It consists of the Woburn Sands Formation, up to about 35 m in present-day thickness, followed by up to 10 m of Gault Clay. These sands and mudstones began to accumulate about 120 million years ago in a sea-way, trending southwest to northeast, which formed when the sea flooded across this part of England. Before the flooding, the area had been dry land for some 30 million years and some of the previously deposited Jurassic sediments had been locally eroded by freshwater streams.

'Flutes' are a feature of large valley slopes in the north of the area where impermeable glacial till and mudstones of the Ancholme Group promote surface drainage and fluvial erosion. It also appears that flutes are more developed on westward-facing valley slopes that receive more solar radiation (and are therefore warmer) than slopes with other aspects. Daily temperature fluctuations, particularly under permafrost conditions, tend to accelerate the breakdown and transportation of rock grains.

Fluted slopes developed on the west-facing margins of the incised meander belt upstream from Bedford

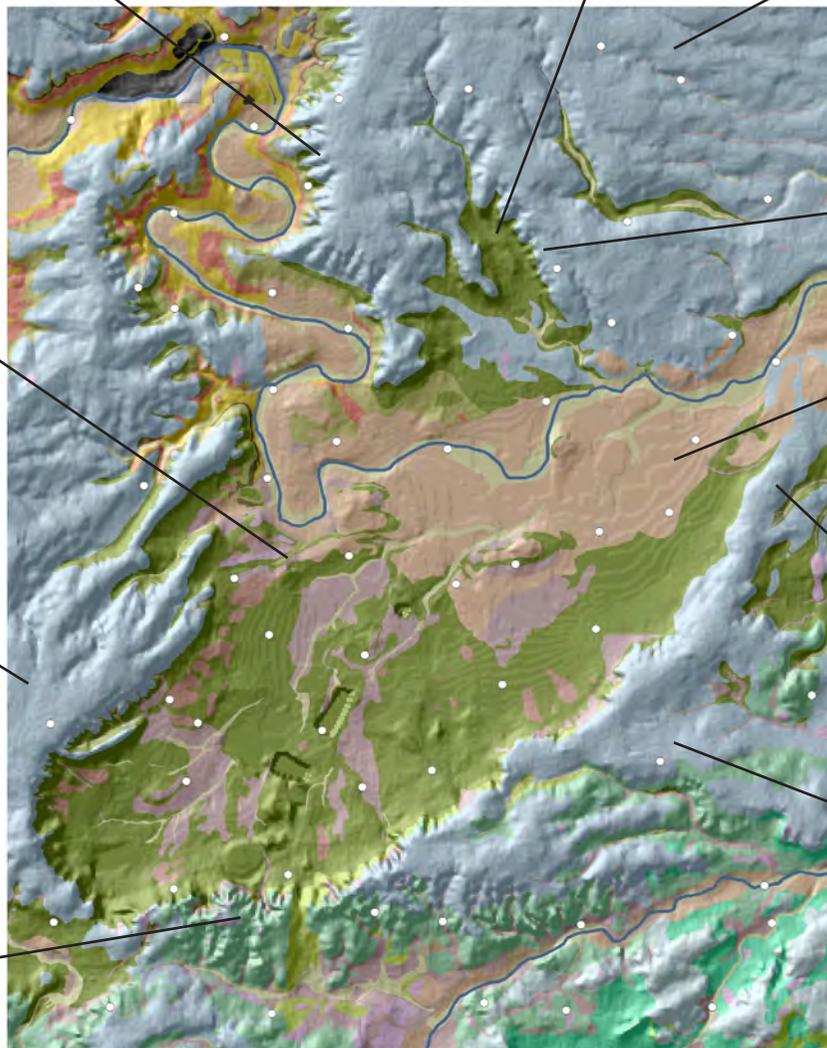
Streams have eroded through the till to reveal the underlying Jurassic mudstones

Upland cover of Anglian till deposited by ice sheets about 450,000 to 400,000 years ago

Downstream from the incised meander belt reaches, the river was free to move laterally and to erode the relatively weak Jurassic mudstones. This lateral shifting, involving large meander loops and perhaps headward erosion by a right bank tributary system, extended the size of the vale to its present form.

Cranfield airfield sited on Anglian till plateau

Fluted margin where resistant sandstones overlie weak Jurassic mudstones to produce a high escarpment (up to c.80 m relief) with local slump features and spring sapped hollows

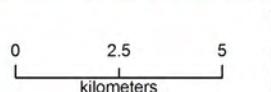


Fluted slopes around Ravensden

Widespread river terrace deposits

Prominent ridge of till with unfluted margins forming the southeast edge of the vale

Upland cover of Anglian glacial till concealing the underlying Woburn Sands Formation



## Map 7. Surface Geology

Almost all of the uplands surrounding Marston Vale are covered by Anglian glacial till. The major drainage valleys and the vale itself have clearly been formed by erosion after the deposition of this glacial till. The presence of 'flutes' along the edges of many of the major slopes provides local evidence of slumping and stream valley erosion during the last few hundred thousand years, and yet these features do not occur everywhere.

Other superficial deposits and associated landforms include buried channels, river terraces, alluvium and head, but these are usually too subtle to be detected using the digital elevation data available in our study.

Marston Vale was eroded by the River Great Ouse as it cut downwards through the glacial till and the underlying bedrock. Where the river system became incised, particularly in the resistant Great Oolite limestones, it was constrained within a narrow meander belt. In contrast, the stretch between Biddenham and Great Barford is relatively straight and contains widespread terrace deposits.

Whilst there is no evidence that moving ice reached this area again during the post-Anglian period, there were certainly prolonged periods during which permafrost conditions prevailed. Modern analogues in the Arctic regions show that under these conditions the varying amounts of rain and snowfall, combined with repetitive seasonal or daily freeze-thaw cycles, greatly accelerate downslope sediment movement. This is particularly pronounced on slopes where gravity promotes slumping and sliding and where there is little or no vegetation to bind the surface layers together. Sediment made available in this way throughout the catchment area is then readily transported by fluvial systems of all scales towards the sea.

River alluvium - deep, stoneless, mainly calcareous clayey soils affected by groundwater

Hanslope soils are heavy calcareous clays which develop on the chalky glacial till

Evesham soils form from the Jurassic clays in depressions where the till has been eroded

Middle Jurassic clay and limestone give rise to well drained calcareous clayey and fine loamy Moreton soils. These can be shallow and brashy in places.

Evesham soils, which are slowly permeable, calcareous clayey, and fine loamy over clayey soils, have developed from the Oxford Clay

Slowly permeable seasonally waterlogged clayey soils of the Denchworth Association form over the Jurassic clays

Deep well drained sandy and coarse loamy soils have developed over the Lower Cretaceous sandstones



The river terraces form wider stretches of deep fine-grained, loamy soils over gravel

Elevated resistant Anglian chalky till gives rise to deep, slowly permeable calcareous clayey, Hanslope soils

The chalky till gives rise to loamy over clayey soils with slowly permeable subsoils and slight seasonal waterlogging

## Map 8. Soils

Much of this area is dominated by the heavy calcareous clay soils of the Hanslope and Evesham Soil Associations. While the Hanslope soils have developed over the clay-rich, calcareous till, the Evesham soils have developed from the Jurassic clays and both result in broadly similar soils. These soils drain slowly and are suitable for winter cereals in drier areas and grasslands where it is slightly wetter.

Because of the areal dominance of these heavy clay soils in Bedfordshire, where differences do occur, local place names often reflect the different geomorphology and soil type. For example, sandy soils occur around Sandy and Chicksands, while Ampthill and Ridgmont are located on elevated areas south of the vale.

Other different soils are found in the lower parts of the river systems. To the north, the shallow and stoney Moreton soils are found on the Great Oolite Group.

On the river terraces, freely draining soils have developed over gravels. Alluvial soils, which have developed on the floodplains, tend to be stoneless, calcareous clays because they are derived from sediment which has been laid down by the river during flood events.